1 Electron density retrieval from truncated Radio 2 **Occultation GNSS** data

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18 Abstract

This paper summarizes the definition and validation of two complementary new strate-

gies, to invert incomplete Global Navigation Satellite System (GNSS) Radio-Occultation

(RO) ionospheric measurements, such as the ones to be provided by the future EUMET-

SAT Polar System 2nd Generation (EPS-SG). It will provide RO measurements with im-

pact parameter much below the LEO height (817 km): from 500 km down approximately.

 The first presented method to invert truncated RO data is denoted as Abel-VaryChap Hybrid modeling from topside Incomplete GNSS RO data (AVHIRO), based on simple First Principles, very precise, and well-suited for post-processing. And the second method is denoted as Simple Estimation of Electron density profiles from topside Incomplete RO data (SEEIRO), is less precise but yields very fast estimations, suitable for Near Real- Time (NRT) determination. Both techniques will be described and assessed with a set ³⁰ of 546 representative COSMIC/FORMOSAT-3 ROs, with relative errors of 7% and 11% ³¹ for AVHIRO and SEEIRO respectively, with 20 minutes and 15 seconds respectively of computational time per occultation in our Intel I7 PC.

³³ 1 Introduction

³⁴ The increase of GNSS RO measurements, such as GPS/MET (Hernández-Pajares, Juan, & Sanz, 2000) CHAMP (Jakowski et al., 2002), COSMIC/FORMOSAT-3 (Olivares-³⁶ Pulido, Hernández-Pajares, Aragón-Angel, & García-Rigo, 2016), GRACE, EPS-SG (Hernández- Pajares et al., 2017), PAZ (Cardellach et al., 2019) and FY3C/GNOS RO (Mao et al., 2016), has allowed for new developments in ionospheric sounding, disposing so far of the complete set of RO measurements, i.e. those with negative elevations referred to the LEO horizon.

 However, in this paper we deal with an initial limitation affecting some missions presently under preparation like EPS-SG: the lack of the topside part of RO observa- tions. A related situation, but without the truncation of RO measurements, happens in the electron density reconstruction from radio occultation data measured by lower al-⁴⁵ titude LEO, e.g. CHAMP. To overcome the upper boundary problem, the inversion as- sisted by an adaptive electron density model of the topside ionosphere and plasmasphere was proposed (Jakowski et al. 2002). But the discontinuity at the transition height should be treated with care. While for truncated radio occultation data we have a double challenge: (1) the missing observations for a significant part of the RO measurements (more than 40% for EPS-SG), and (2) the longer length through the blind area makes the re- trieval results more sensitive to the accuracy of the model. This last point has required the implementation of an still more realistic topside electron density model, the linear Vary-Chapman one, taking advantage of recent results. We showed in (Hernández-Pajares et al., 2017) that once the electron density profile is well known below such ceiling height, e.g. 500 km, it can be easily and accurately extrapolated. In this manuscript we will fo- cus on the pending previous problem, that is, how to determine the electron density pro- file under such topside-truncated set of dual-frequenty RO GNSS measurements, gath- ered from the LEO height, e.g. approximately 800 km, paying attention to the accuraccy but also to the computational load.

 Indeed, the dual-frequency measurements provided by GNSS receivers on board Low Earth Orbiters (LEO) in RO scenarios, with negative elevation angle, are very sen- sitive to vertical variability. This fact allows the retrieval of the electron density profiles below the LEO height. The new EPS-SG satellites at 817 km height are conceived for neutral atmospheric sounding. Nevertheless this provides also opportunities for ionospheric sounding, but with RO measurements only taken with impact parameter height below 500 km (see layout in Figure 1). Some aspects of the electron density retrieval and im- pact on EPS-SG have been already studied, in particular a new electron density Vary-Chapman Extrapolation Technique (VCET) for impact parameters of 500km up to the

Figure 1. Layout of the incomplete RO measurements scenario studied in this work, showing up, as conceptual example, some transmitter-receiver rays illuminating three layers, in green, magenta and orange colors. The height interval used to fit the Vary-Chap model is also represented (light blue color).

 EPS-SG orbital height (Hern´andez-Pajares et al., 2017). This model can be considered σ an improvement of previous extrapolation approaches, e.g. (Jakowski, 2005), which as- π sumed a fixed scale height Chapman model for ionosphere in combination with an ex- ponential model for plasmasphere. VCET is based on the predominant linear increase observed above hmF2 of scale height, due to its proportional dependence on tempera- ture, as it was shown in (Olivares-Pulido et al., 2016), and in the recent climatic mod- els (Prol, Hern´andez-Pajares, Camargo, & Muella, 2018) that are based as well on top- sounding data (Prol, Themens, Hern´andez-Pajares, Camargo, & Muella, 2019). VCET assumes the availability of a properly estimated electron density profile.

 Nevertheless before applying such extrapolation, there is the need of an accurate ⁷⁹ estimate of the electron density profile below the first impact parameter height with mea- surements (500 km). In this regard we introduce two new techniques: the Abel-VaryChap Hybrid, and the Simple Estimation of Electron density profiles, both modeling from top-⁸² side Incomplete RO data (AVHIRO and SEEIRO, respectively). The two approaches do not depend on external models or data beyond the radio-occultation measurements. AVHIRO priorizes the accuracy and SEEIRO priorizes the estimation computating time in order to allow near real-time usage. The assessment of both techniques is presented with mea- surements of COSMIC/FORMOSAT-3 truncated up to 500 km, during 4 representative ⁸⁷ periods, by comparing the electron density profiles with the corresponding ones obtained from the full radio-occultation measurements.

⁸⁹ The starting point for both approaches is the set of radio-occultation measurements ⁹⁰ gathered from the LEO, in our case the COSMIC/FORMOSAT-3 at a height of about $r_L = 800$ km. And they are truncated to a maximum impact parameter height equal ϵ_{92} to the expected value of $r_0 = 500$ km for EPS-SG. The dual-frequency ionospheric phase 93 combinations, $L_I = L_1 - L_2$, are corrected by substracting the Slant Total Electron ⁹⁴ Content (STEC) above the LEO orbit. These STEC values are computed from the Pre-⁹⁵ cise Orbit Determination (POD) antenna measurements, i.e. with positive elevations, ⁹⁶ by means of a dual-layer tomographic voxel model which simultaneously estimate the ⁹⁷ electron density of the topside voxels and the carrier phase ambiguities, as described in 98 (Hernández-Pajares et al., 2017).

⁹⁹ 2 Abel-VaryChap Hybrid modeling from topside Incomplete Radio ¹⁰⁰ Occultation data (AVHIRO)

The first method that we introduce in this paper is the hybrid approach- AVHIRO, which consists of Abel inversion and Vary-Chap model. It synergistically solves the full electron density, ambiguity term and four parameters of Vary-Chap model at the same time, taking into account the nonlinear interactions between the unknown parameters. As it has been indicated above, Vary-Chap model summarizes the expected distribution of the topside electron density. And it is applied as constraint to improve the accuracy of the overall electron density estimation. Specifically, and following Figure 1, we can relate the ionospheric combination in length units, $L_I = L_1 - L_2$, where L_1 and L_2 are the carrier phases measured in frequencies f_1 and f_2 , with the known crossing lengths $l_{j,k}$ of the corresponding j-th line-of-sight with each given k-th layer, and with the unknown electron density values N and carrier phase ambiguity in length units B_I . Following (Hernández-Pajares et al., 2011), B_I contains the integer terms in cycles, $\lambda_m N_m$, and instrumental phase delays for receiver and satellite, respectively δb_m , $\delta b'_m$, and for frequencies $m = 1, 2$:

$$
B_1 = B_1 - B_2 = \lambda_1 N_1 - \lambda_2 N_2 + \delta b_1 - \delta b_2 + \delta b_1' - \delta b_2'
$$

¹⁰¹ In this context we can express the ionospheric combination of carrier phases, from top ¹⁰² to bottom, as:

$$
(L_I)_1 = \alpha (2l_{1,1}N_1 + 2l_{1,2}N_2 + \dots + 2l_{1,x}N_x) + B_I
$$

\n
$$
(L_I)_2 = \alpha (2l_{2,1}N_1 + 2l_{2,2}N_2 + \dots + 2l_{2,x}N_x + 2l_{2,x+1}N_{x+1}) + B_I
$$

\n
$$
(L_I)_3 = \alpha (2l_{3,1}N_1 + 2l_{3,2}N_2 + \dots + 2l_{3,x}N_x + 2l_{3,x+1}N_{x+1}) + B_I
$$

\n
$$
\dots
$$

\n
$$
(L_I)_6 = \alpha (2l_{6,1}N_1 + 2l_{6,2}N_2 + \dots + 2l_{6,x}N_x + 2l_{6,x+1}N_{x+1} + 2l_{6,x+2}N_{x+2}) + B_I
$$

\n
$$
\dots
$$

\n(1)

103 The observation set of equations 1 can be summarized in matrix notation as $Ax =$ $b, \text{ where } x = (N_1, \ldots, N_x, \ldots, B_I)^T \text{ and } A_{j,k} = 2\alpha \cdot l_{j,k} \text{ for } k < M \text{ and } A_{j,M} = 1$ for the ambiguity coefficient, being $\alpha = 1.05 \times 10^{-17}$ m³ the scaling factor converting $_{106}$ electron content into delay (see (Hernández-Pajares et al., 2011)) and being M the num-¹⁰⁷ ber of unknowns.

 From these equations we cannot apply directly the Abel inversion algorithm (see for instance (Hernández-Pajares et al., 2000)), because the design matrix \vec{A} is rank de- ficient due to the lack of observations above 500 km. In order to solve such rank-defect equations, the Vary-Chap model is added as constraint above 500 km, as it was indicated above, with parameters mainly estimated within the height range of 380 km to 430 km. The Vary-Chap model is based on the physics of the problem as it has been mentioned above, allowing an extrapolation compatible with the observational data.

 The Vary-Chap model is based on a non-linear interaction between the parame- ters to be estimated. So in order to use it as prior, it is important to be aware of the dif- ficulties related to the parameter estimation. The structure of the Vary-Chap model con-118 sists of two exponential terms of a variable z, as can be seen in Eq.(2). This variable z consists of terms that have to be estimated such as h_m and H. The estimation of these parameters by means of the conventional methods based on gradient search is difficult 1_{121} (Luenberger, Ye, et al., 1984). This difficulty is due to the fact that small variations in the parameters give rise to very large changes in the value of the derivative, caused by the dependencies inside the exponential functions. These large changes are explained by the different scale of the values of the parametes and also by the nonlinear terms of the expression, which can be summarized as a multiplicative interaction between the paramters (i.e. N_m), the exponencial of the inverse of a variable (i.e. $e^{\frac{1}{H}}$) and a double ex- $_{127}$ ponential of z.

 A family of estimation methods that is robust to the problems of differentiability of the target function are the algorithms for minimization without derivatives. This fam- ily of algorithm search for the optimum by comparing perturbations of a given candi- date to the solution. The fact that instead of computing a derivative, they rely on com- parisons, removes the problems related to the extreme variability of the derivatives, and the different scale of the unknowns. There are different algorithms in this family such as the Powell's method, the Nelder Mead algorithm and the pattern search, see for in- stance (Press, Flannery, Teukolsky, Vetterling, et al., 1989). Due to the differences of scale of the parameters to be estimated, we selected the algorithm that in principle is more rubust to this phenomenon, which is the Powell search method (Powell, 1964). This al- gorihm was used in order to estimate the electron densities below 380 km, the ambiguity term and four Vary-Chap parameters, which are peak height h_m , peak eletron den-140 sity N_m , scale height at peak H_0 and the derivative of scale height $\partial H/\partial h$, simultane- ously. Afterwards, the electron densities above 380 km are updated from the Vary-Chap model using the estimated parameters. The expression that relates the electron density N , with the height h, is the following,

$$
N = N_m e^{\frac{1}{2}(1 - z - e^{-z})}, \quad where \ z = \frac{h - h_m}{H}
$$
 (2)

145

$$
H = \frac{\partial H}{\partial h}(h - h_m) + H_0 \tag{3}
$$

¹⁴⁷ where N and H represent the electron density and scale height at height h above peak ¹⁴⁸ respectively.

The cost function of powell search is mainly composed of two terms: one is $||Ax - b||^2$ 149 150 and the other is the difference with respect to a reference estimate x_0 , weighted by λ which 151 is a regularization parameter, $\lambda \|x - x_0\|^2$. The regularization parameter λ controls the 152 smoothness of the estimation. Besides, additional penalization terms on h_m , N_m , H_0 and $\partial H/\partial h$ are added to the cost function to constrain Vary-Chap parameters in a realis-¹⁵⁴ tic range.

¹⁵⁵ The unknown vector x in equation $Ax = b$ is composed to three parts x^1 (electron densities from 380 km and 1000 km), x^2 (electron densities below 380 km) and x^{ambi} . ¹⁵⁷ The iterative algorithm for the estimation is described as follows:

- ¹⁵⁸ 1. Initial electron density profile below 500 km and ambiguity term are derived from ¹⁵⁹ Abel inversion and then the full profile is extrapolated to 1000 km. These values 160 define the vector x_0 . Next we iterate updating this vector $x_0 = [x_0^1, x_0^2, x_0^{ambi}]$ ¹⁶¹ each time, following next point.
- ¹⁶² 2. The terms x_0^2 (current estimate of the electron densities below 380 km) and x_0^{ambi} , are extracted from current unknowns x_0 , together with x_0^1 , which is calculated with ¹⁶⁴ the Vary-Chap model's parameters $[h_m, N_m, H_0, \partial H/\partial h]_0$. These parameters are 165 initialized at the beginning with typical values: h_m and N_m are derived from the ¹⁶⁶ very first Abel inversion neglecting the electron content above 500 km, $H_0 = 30$ km and $\partial H/\partial h = 0.05$ from (Olivares-Pulido et al., 2016). They form the x_{powell} .
- ¹⁶⁸ 3. The cost function of the Powell search is the sum of $||Ax_{powell} b||^2$ and $\lambda ||x_{powell} x_0||^2$, 169 where λ is obtained from the ratio of both estimated standard deviations in the ¹⁷⁰ previous iteration: the one from post-fit residuals vs the one for peak electron den-¹⁷¹ sity. And the Powell algorithm is applied in order to estimate un update of the v_1 ₁₇₂ values $[x^2, x^{ambi}, h_m, N_m, H_0, \partial H/\partial h].$
- ¹⁷³ 4. The solution at the ith iteration x_{powell_i} can be computed by the searched vec-¹⁷⁴ tor $[x^2, x^{ambi}, h_m, N_m, H_0, \partial H/\partial h]_i$. Then we update x_0 by assigning x_{powell_i} to x_0 and we go to step 2. Empirically 10 iterations are performed in order to en-¹⁷⁶ sure the convergence of the algorithm.

¹⁷⁷ 3 Simple Estimation of Electron density profiles from topside Incom-¹⁷⁸ plete Radio Occultation data (SEEIRO)

 This is the second method that we introduce in this paper which, in contrast with the preceeding one, trades accuracy with speed. The SEEIRO algorithm, iteratively es-¹⁸¹ timates the electron density profiles under the assumption of an exponential behaviour of the electron density among consecutive values in height. In this way, one variable scale height per topside height below 500 km can be easily obtained without the knowledge of h_m and N_m , and the Linear Vary-Chap model can be fitted from them and used for extrapolation, and correcting the L_I observations for next iteration.

¹⁸⁶ In the first iteration, $i = 0$, the system is initialized using only the Abel inversion ¹⁸⁷ of the available measurements below $r_0 = 500 \text{km}$, and neglecting the electron density ¹⁸⁸ for higher altitudes, $r_0 < r < r_L$, and estimating simultaneously the carrier phase ambiguity, B_I^0 and the electron density $N_e^0(r_k)$, for $k = 1, ..., M$, being M the number 190 of layers defined with a width of Δr (e.g. $\Delta r = 3 \text{km}$).

 1_{91} In the subsequent iterations i, we focus now on the top values of the previous $(i-$ ¹⁹² 1) solution, above the F_2 peak geocentric distance r_{mF2} and below the highest available geocentric distance r_0 and with a tolerance ϵ : $N_e^{(i-1)}(r_k)$ for $r_{mF2} + \epsilon \leq r \leq r_0 - \epsilon$. ¹⁹⁴ From these values we estimate the scale height, assumed constant for each pair of con-¹⁹⁵ secutive values only. Indeed, we can approximate the Chapman function, equation 4, by 196 the exponential approximation specially valid when $z \gg 1$:

$$
N = N_m e^{\frac{1}{2}(1-z)}\tag{4}
$$

¹⁹⁸ Then we can obtain the corresponding scale height values without the dependence on 199 the F_2 peak height and density values, from two consecutive values:

$$
H(r_k) \simeq \frac{-\Delta r}{2 \cdot \ln \frac{N_e^{(i-1)}(r_k)}{N_e^{(i-1)}(r_k - \Delta r)}}\tag{5}
$$

²⁰¹ From the series of scale height values $H(r_k)$ the linear fit is performed following ²⁰² equation 3 removing iteratively outliers with residual greater than 2.5 times its standard ²⁰³ deviation.

²⁰⁴ From the resulting linear model, the scale height is extrapolated, and a constant $_{205}$ value H_0 is adopted when the estimated vertical gradient is not positive, i.e. approxi-²⁰⁶ mately in the 10% of cases. Afterwards, the electron density is consistently obtained for $207 \, r > r_0$, with the exponential approximation equivalent to equation 5 i.e.:

$$
N_e(r) = N_e(r - \Delta r) \cdot e^{-\frac{\Delta r}{2\left[H_0 + \frac{\partial H}{\partial h}(h - h_m)\right]}} \tag{6}
$$

 \mathbb{R}^2 From these values, in the given iteration, the STEC between r_0 and r_L is mitigated $_{210}$ within the measurements with impact factor below r_0 , and a new Abel inversion is per-²¹¹ formed repeating the same procedure described above, up to 10 times. This number of ²¹² iterations is an optimal value empirically obtained, which ensures the convergence.

²¹³ 4 Estimation assessment

 In order to assess the performance of both AVHIRO and SEEIRO in the height range corresponding to the observations impact parameter heights (below 500 km), we have considered the selected set of 570 radio-occultations corresponding to the first day of each one of the four weeks studied in (Hern´andez-Pajares et al., 2017). They are represen-tative of the previous solar cycle (see Figure 2).

Figure 2. Solar Flux, Kp index during the four selected periods, extracted from (Hernández-Pajares et al., 2017).

 One first illustrative example is shown in Figure 3. The performance of one typ- ical occultation retrieval is compared between not applying (initial values) and apply- $_{221}$ ing these new techniques (last iteration). It can be seen that, in this case, the error goes down from 45-46%, to 1.3% with AVHIRO and 10.0% with SEEIRO. We will see below that these final relative errors are not far from the most frequent ones.

 We have considered the first day of each one of the four representative periods: Namely day 346 of year 2006 (low solar flux, before a major geomagnetic storm), day 234 of year 2008 (low-mid solar flux), and days 261 and 352 of year 2011 (high solar flux). The com- parison of the absolute and relative error RMS for AVHIRO are respectively shown in F_{228} Figures 4 and 5. The error reduces from $1.0 \times 10^{11} \pm 1.3 \times 10^{11} \text{m}^{-3}$ (51.6% of RMS) in the initial iteration to $1.5 \times 10^{10} \pm 2.6 \times 10^{10} \text{m}^{-3}$ (13.1% of RMS) in the final one.

230 The histogram peak, i.e. the mode, of relative error is 3% for AVHIRO versus 6- 10% for SEEIRO (see Figure 6). Once we remove the values with relative error higher than 20% (23 of 570 radio-occultations, i.e. the 4% of values) the relative error decreases to 7.2%, clearly below the corresponding value for SEEIRO: 10.6% (see again Figure 6). Nevertheless SEEIRO is 70 times faster than AVHIRO, with an average processing time per radio occultation in our Linux I7 PC of 15 seconds, vs 20 minutes with AVHIRO.

 These results strongly suggest that SEEIRO and AVHIRO techniques are appropi- ate, respectively, for near real-time and postprocessing determination of electron den- sity profiles from topside-truncated radio occultation data. A comparison of the main characteristics of both techniques is summarized in Table 1.

5 Extrapolation assessment

 Although the area below 500 km, tackled in previous section, is the main target of this work, the extrapolation precision for the topside part (above 500 km in this work) should be examined, since the electron densities in the blind area have a non-negligible impact on the retrieval. In AVHIRO method, the full electron densities are estimated

Figure 3. Example of the electron density (blue points) obtained from the measurements below 500 km of impact parameter height, with the AVHIRO (top row) and SEEIRO (bottom row) approaches, comparing the first and last iteration included in left- and right-hand columns respectively. It corresponds to a single radio-occultation, of satellite PRN13 with measurements from COSMIC/FORMOSAT-3 receiver L261 starting on second 37323 of day 261 of year 2011. They are compared with two different solutions obtained from the complete set of measurements. The first one has been obtained by applying Abel inversion under the assumption of spherical Symmetry (green points) and the second one modelling the horizontal variability with the Separability concept mentioned above; the profile corresponding to hmF2 tangent point is represented with red points. In both reference cases POD-data based LEO topside corrections have been applied.

 simultaneously, with a full linear Vary-Chap model for the topside part of the electron density profile, instead of by two steps to separate the observed and blind area. Hence ²⁴⁸ the topside assessment can to some extent reflect the performance of Vary-Chap model, which is shown in this section for completeness.

 From absolute errors histogram (see Figure 7), the performance in the blind area ²⁵¹ is a little worse than but comparable to that in the area below 500 km, with bias $2.0\times$ 10^{10} m⁻³ and standard deviation 4.0×10^{10} m⁻³, which are in the same magnitude as those in the lower part. While the relative errors (see Figure 8) are quite large with 53.3% of RMS compared to 13.1% in the observed area. This can be easily explained by tak- ing into account that the electron densities above 500 km are quite small and the sam- ple number is limited for statistics, so the small absolute errors of electron density could produce big relative errors. For example, the four cases in Figure 9 show that, even with high relative error (93.1%), the extrapolation results by Vary-Chap model are very close to reference values, with 8.9×10^9 m⁻³ of absolute error inside error bars. Therefore,

Figure 4. Histogram of the electron density error RMS values, one per occultation and expressed in m⁻³. They correspond to the initial (left-) and final (right-hand plot) iteration of AVHIRO, for the selected COSMIC/FORMOSAT-3 radio-occultations during days 346 of year 2006, 234 of year 2008, and days 261 and 352 of year 2011.

Figure 5. Histogram of the relative electron density error RMS values, one value per occultation and expressed in %. They correspond to the initial (left-) and final (right-hand plot) iteration of AVHIRO, for the selected COSMIC/FORMOSAT-3 radio-occultations during days 346 of year 2006, 234 of year 2008, and days 261 and 352 of year 2011.

Figure 6. Histogram of the relative electron density error RMS values, one per occultation and expressed in %, focused on the values smaller than 20%, the 96% of the analyzed radiooccultations. They correspond to the final iterations of AVHIRO (left-hand plot) and of SEEIRO (right-hand plot), for the selected COSMIC/FORMOSAT-3 radio-occultations during days 346 of year 2006, 234 of year 2008, and days 261 and 352 of year 2011.

	AVHIRO	SEEIRO
Ne Relative Accuracy	7%	11%
Predominant Ne Rel. Acc.	3%	$6 - 10\%$
Ne Absolute Accuracy	$(1.5 \pm 2.6) \times 10^{10} \text{m}^{-3}$	$(3.9 \pm 2.3) \times 10^{10}$ m ⁻³
Predominant Ne Abs. Acc.	$< 10^{10}$ m ⁻³	10^{10} m ⁻³
CPU time per preprocessed RO	20 minutes	15 seconds
Suitable for NRT service?	Not now	Yes
Required ancillary information?	$\rm No$	No
Required inputs	2-freq. GPS carrier phase meas.,	
	predicted GPS and LEO orbits (both)	
Convenient inputs	2-freq. GPS POD carrier phase meas. (both)	

Table 1. Pros and Cons of AVHIRO vs SEEIRO: Summary

Figure 7. Histogram of the absolute topside electron density error RMS values, one value per occultation and expressed in m^{-3} . They correspond to the final iteration of AVHIRO, for the selected COSMIC/FORMOSAT-3 radio-occultations during days 346 of year 2006, 234 of year 2008, and days 261 and 352 of year 2011.

²⁶⁰ this proves good performance of Vary-Chap model in simultaneously extrapolating the ²⁶¹ topside electron density as well.

²⁶² 6 Conclusions

 In this work we have presented a new Abel-VaryChap Hybrid modeling from top- side Incomplete RO data (AVHIRO). This can complete the set of algorithms for iono- spheric electron density retrieval from GPS RO data in EPS-SG with lack of measure-ments for impact parameter heights above 500 km, as a new post-processing technique.

²⁶⁷ AVHIRO reduces, without the need of external data, the electron density error of ₂₆₈ the RO inversion with measurements up to 500 km regarding to the full inversion with b observations up to 800 km: from 51.6% before, to 13.1% percentage of electron density ²⁷⁰ RMS after applying AVHIRO. In particular it reduces the predominant relative error to ²⁷¹ 3% compared with the 6-10% obtained with the fast Simple Estimation of Electron den-

Figure 8. Histogram of the relative topside electron density error RMS values, one value per occultation and expressed in %. They correspond to the final iteration of AVHIRO, for the selected COSMIC/FORMOSAT-3 radio-occultations during days 346 of year 2006, 234 of year 2008, and days 261 and 352 of year 2011.

Figure 9. Four representative cases showing the EDP obtained with AVHIRO applied to the COSMIC/FORMOSAT-3 measurements below 500 km, compared with the EDP obtained from the full RO dataset.

sity profiles from topside Incomplete RO data (SEEIRO) approach. Moreover AVHIRO

provides simultaneously the linear Vary-Chapman extrapolated electron density profile

 with accuracies just slightly lower than those obtained at heights below 500 km with observations: $(2.0 \pm 4.0) \times 10^{10} \text{ m}^{-3}$ above vs $(1.5 \pm 2.6) \times 10^{10} \text{ m}^{-3}$ below 500 km.

 Further potential improvements of the technique can be studied in future works, ₂₇₇ in particular to try to speed up AVHIRO to be hopefully suitable for NRT.

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used in this work is available at ftp://chapman.upc.es/.AVHIRO and SEEIRO results

or alternatively it can be requested to Manuel Hern´andez-Pajares (manuel.hernandez@upc.edu).

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Figure 1.

GPS

Figure 2.

 \top $\frac{1}{2}$ 2 \leq $10 - 22$ $\overbrace{}$ Flux Solar

Solar Flux and KP indices of proposed scenarios

Figure 3.

RO-11_261_l220_13_37323_6, AVHIRO rel.error: 45.1 (iter.#0)

Figure 4.

Absolute_error_Ne_for_50km_to_500km

AVHIRO abs.error (iter.#9)

Figure 5.

Figure 6.

SEERO rel.error (iter.#9, remove outliers greater than 20%)

Figure 7.

AVHIRO blind area rel.error (iter.#9)

Figure 8.

AVHIRO blind area rel.error (iter.#9)

Figure 9.

